Overview WP1 (reanalysis production) and WP5 (service developments)

Patrick Laloyaux & Manuel Fuentes

















STANDARD

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	Operational Operational datasets are the forecasts output by our current model. - Medium range - Extended Range - Long Range Real-time data Atmospheric HRES 10-day Ocean Wave 10-day Atmospheric ENS 15-day Ocean Wave ENS-WAM 15-day Atmospheric ENS ext 46-day	Climate reanalysis ECMWF uses its forecast models and data assimilation systems to 'reanalyse' archived observations, creating global data sets describing the recent history of the atmosphere, land surface, and oceans. Browse reanalysis datasets	Atmospheric composition Datasets for atmospheric composition from the Copernicus Atmosphere Monitoring Service (CAMS) combine atmospheric modelling with Earth observation data to provide information covering European air quality, global atmospheric composition, climate forcing, the ozone layer, UV and solar energy, and emissions and surface fluxes. CAMS forecasts	One access point for the reanalysis datasets										

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User statistics for CERA-20C

- Advertised on the main reanalysis websites
- E-mail sent to all the ERA-20C users
- ECMWF newsletter published
- Datasets presented in many conferences





How much coupling in the CERA system?

Patrick Laloyaux, Sergey Frolov (NRL) and Massimo Bonavita







Outer loop coupling in the CERA system

Iterative process where the ocean and the atmosphere converge towards a consistent coupled state

 $\begin{bmatrix} \mathbf{x}^{\mathbf{0}} \\ \mathbf{x}^{\mathbf{0}} \end{bmatrix} = \begin{bmatrix} \mathbf{x}^{\mathbf{b}} \\ \mathbf{x}^{\mathbf{b}} \end{bmatrix}$

for k=0,1,... do

Compute observation departures

$$\begin{bmatrix} \delta \mathbf{y}^{\mathbf{k}} \\ \delta \mathbf{y}^{\mathbf{k}} \end{bmatrix} = \begin{bmatrix} \mathbf{y} \\ \mathbf{y} \end{bmatrix} - \begin{bmatrix} \mathcal{H} \\ \mathcal{H} \end{bmatrix} \mathcal{M}(\mathbf{x}^{\mathbf{k}}, \mathbf{x}^{\mathbf{k}})$$

Compute increments

$$\delta \mathbf{x}^{\mathbf{k}} = (\mathbf{x}^{\mathbf{b}} - \mathbf{x}^{\mathbf{k}}) + \mathbf{B}\mathbf{H}^{\mathrm{T}}(\mathbf{H}\mathbf{B}\mathbf{H}^{\mathrm{T}} + \mathbf{R})^{-1}\delta \mathbf{y}^{\mathbf{k}}$$

 $\delta \mathbf{x^k} = (\mathbf{x^b} - \mathbf{x^k}) + \mathbf{B}\mathbf{H}^{\mathrm{T}}(\mathbf{H}\mathbf{B}\mathbf{H}^{\mathrm{T}} + \mathbf{R})^{-1} \delta \mathbf{y^k}$

Update initial condition

$$\begin{bmatrix} \mathbf{x}^{k+1} \\ \mathbf{x}^{k+1} \end{bmatrix} = \begin{bmatrix} \mathbf{x}^{k} \\ \mathbf{x}^{k} \end{bmatrix} + \begin{bmatrix} \delta \mathbf{x}^{k} \\ \delta \mathbf{x}^{k} \end{bmatrix}$$
end

Outer loop coupling in the CERA system

Iterative process where the ocean and the atmosphere converge towards a consistent coupled state

 $\begin{bmatrix} \mathbf{x^0} \\ \mathbf{x^0} \end{bmatrix} = \begin{bmatrix} \mathbf{x^b} \\ \mathbf{x^b} \end{bmatrix}$

for k=0,1,... do

 $\begin{array}{l} \text{Compute increments} \\ \delta \mathbf{x}^{k} = (\mathbf{x}^{b} - \mathbf{x}^{k}) + \mathbf{B}\mathbf{H}^{T}(\mathbf{H}\mathbf{B}\mathbf{H}^{T} + \mathbf{R})^{-1} \left(\mathbf{y} - \mathcal{H}\mathcal{M}(\mathbf{x}^{k}, \mathbf{x}^{k})\right) \\ \delta \mathbf{x}^{k} = (\mathbf{x}^{b} - \mathbf{x}^{k}) + \mathbf{B}\mathbf{H}^{T}(\mathbf{H}\mathbf{B}\mathbf{H}^{T} + \mathbf{R})^{-1} \left(\mathbf{y} - \mathcal{H}\mathcal{M}(\mathbf{x}^{k}, \mathbf{x}^{k})\right) \\ \text{Update initial condition} \\ \begin{bmatrix} \mathbf{x}^{k+1} \\ \mathbf{x}^{k+1} \end{bmatrix} = \begin{bmatrix} \mathbf{x}^{k} \\ \mathbf{x}^{k} \end{bmatrix} + \begin{bmatrix} \delta \mathbf{x}^{k} \\ \delta \mathbf{x}^{k} \end{bmatrix} \\ \text{end} \end{array}$

Atmospheric increment depends on the current ocean state Ocean increment depends on the current atmospheric state

Single observation experiments

ATMOS. BOUNDARY LAYER Compare the implicit cross-correlations from CERA with explicit cross-correlation from an ensemble

Single observation experiments

- the true state is known
- one accurate and unbiased observation (T, first level in the ocean, valid after 10 hours)
- background from another run with a larger error

OCEAN BOUNDARY LAYER

How CERA transfers information in space

Coupled state estimate after 10 hours



How CERA transfers information in space

Coupled state estimate after 10 hours



How CERA transfers information in space

Coupled state estimate after 10 hours



How CERA transfers information in time

Ocean and atmospheric trajectories during the CERA minimisation



How CERA transfers information in time

Ocean and atmospheric trajectories during the CERA minimisation



How CERA transfers information in time

Ocean and atmospheric trajectories during the CERA minimisation



Explicit cross-correlations from an ensemble

 $\begin{vmatrix} \mathbf{x}^{\mathbf{f}} \\ \mathbf{x}^{\mathbf{f}} \end{vmatrix} = \begin{vmatrix} \mathbf{x}^{\mathbf{b}} \\ \mathbf{x}^{\mathbf{b}} \end{vmatrix}$

Compute observation departures

$$\begin{bmatrix} \delta \mathbf{y} \\ \delta \mathbf{y} \end{bmatrix} = \begin{bmatrix} \mathbf{y} \\ \mathbf{y} \end{bmatrix} - \begin{bmatrix} \mathcal{H} \\ \mathcal{H} \end{bmatrix} \mathcal{M}(\mathbf{x}^{\mathbf{k}}, \mathbf{x}^{\mathbf{k}})$$

Compute analysis

Explicit coupled background error, computed from an ensemble of 25 coupled forecasts.

Used to update the background state following the Kalman filter equation



Explicit cross-correlations from an ensemble

Coupled state estimate after 10 hours



Implicit cross-correlation from CERA and explicit cross-correlation from an ensemble correct the coupled state in a similar way

What's happening when the mixed layer is deep?

Ocean and atmospheric trajectories during the CERA minimisation



What's happening when the mixed layer is deep?

Ocean and atmospheric trajectories during the CERA minimisation



What's happening when the mixed layer is deep?

Ocean and atmospheric trajectories during the CERA minimisation



Why does the 0.3 ocean temperature increment disappear after 2 hours?

Summary plot of the 18 single observation experiments



Summary plot of the 18 single observation experiments

Evolution of the atmospheric increment (normalised by the ocean increment)



Tropics → Most of the ocean increment is transferred to the atmosphere within 10 hours

Extra-tropics (shallow mixed layer) → Most of the ocean increment is transferred to the atmosphere within 10 hours

Extra-tropics (deep mixed layer) → Atmospheric response is extremely noisy